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del Mar**



# ***THEORETICAL MODELING OF DUAL-FREQUENCY SCATTEROMETER RESPONSE: IMPROVING OCEAN WIND AND RAINFALL EFFECTS***

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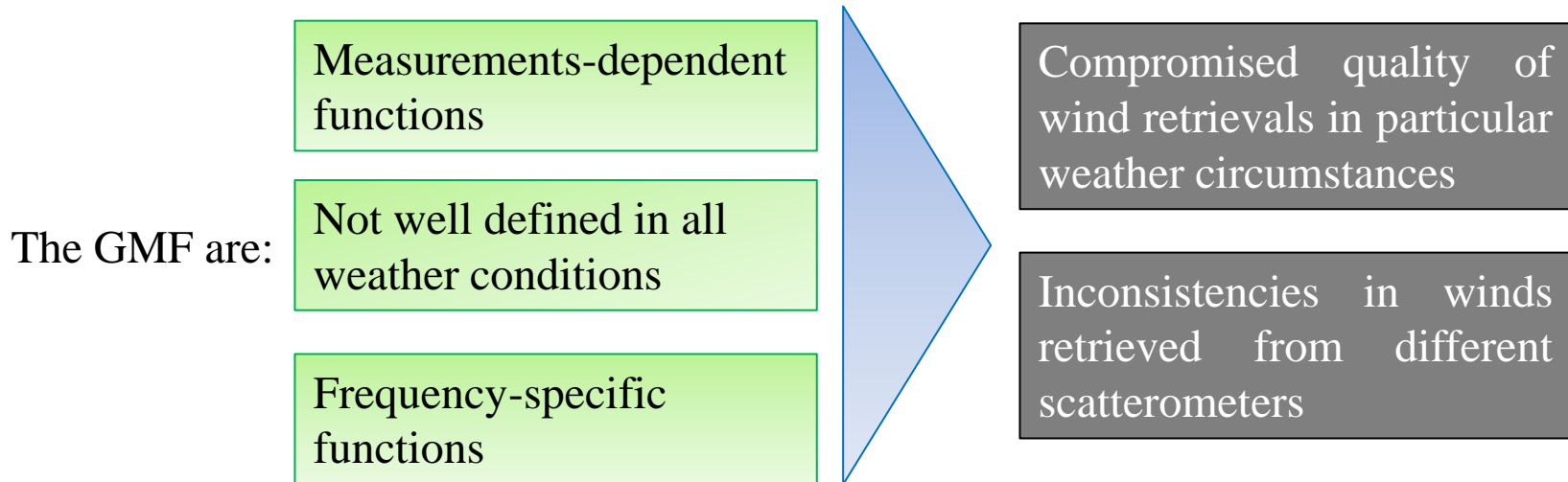
# Outline

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- Theoretical Modeling in Presence of Wind
- Theoretical Modeling in Presence of Both Wind and Rain
- Conclusion

# Introduction

Good confidence in the use of the **Geophysical Model Functions (GMFs)** to define the relationship between the ocean surface backscattering coefficient ( $\sigma^0$ ) and **wind speed (ws)** / **wind direction ( $\phi_R$ )** and to perform wind retrievals. However:



# Goals: Improving Theoretical Models

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The theoretical models have the big potential to:

- describe the **physical-based relation** between the **radar backscattering** and the **ocean surface wind**
- Account for the additional effects compromising the surface backscattering and it turn the wind measurements.

It requires:

- full understanding of the mechanisms regulating the relation between the ocean wind and waves;
- accurate physics-based parameterization of such mechanisms;
- accurate parameterization of the ocean surface itself.

The *objective* is to **improve the theoretical modeling** of the ocean scatterometer backscatter, accounting for the different sources that the current sensor-specific GMFs are not able to characterize, such as **frequency-specific sensor sensitivity to wind, rain effects** and different viewing geometries.

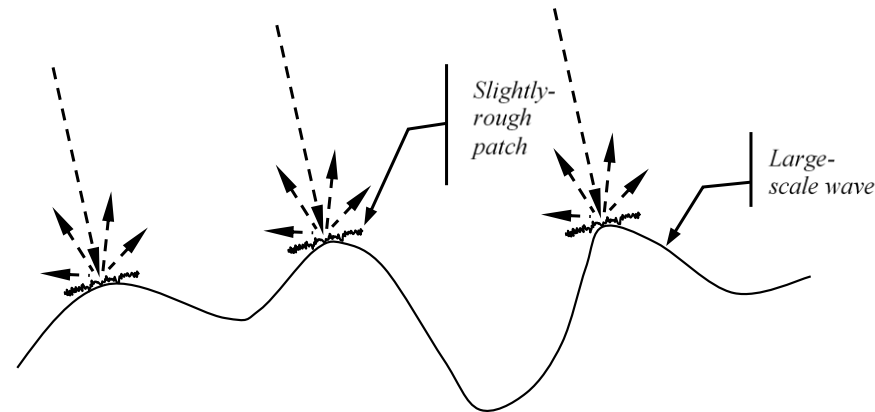
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# Two Scale Model (TSM)

- The ocean surface is assumed as composed by small-scale **wind-driven capillary waves** (slightly-rough patches) superimposed to large-scale **gravity waves**;



- The backscattering coefficient  $\sigma^0$  is defined as the sum of **quasi-specular backscattering coefficient**  $\sigma^{0(0)}$  from the large-scale waves and the contribution of the **small-scale roughness component**  $\sigma^{0(1)}$ :

$$\sigma_{pq}^0(\theta, \varphi) = \sigma_{pq}^{0(0)}(\theta, \varphi) + \sigma_{pq}^{0(1)}(\theta, \varphi)$$

integral of the radiation from the single slightly-rough patch, weighted by the probability density function of the long waves slopes  $P(S_x, S_y)$

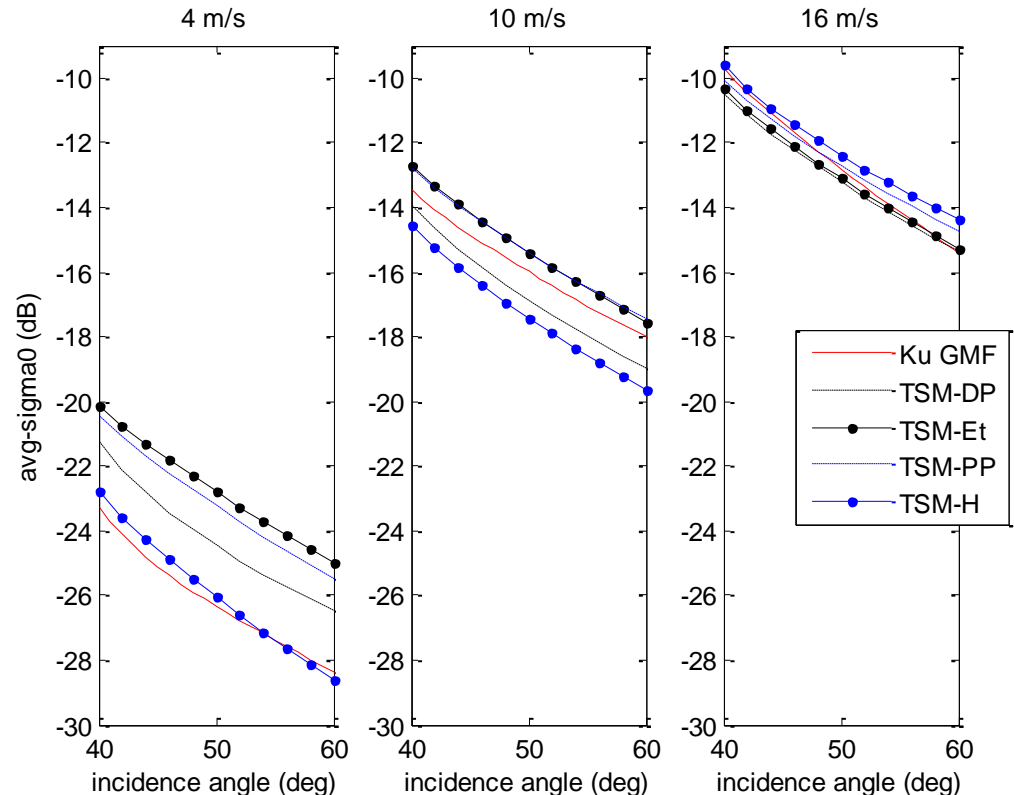
**An accurate representation of the ocean surface roughness is crucial**

# Ocean Wave Spectrum Models (1/2)

The wave spectrum comes from the equation describing the ocean waves evolution, regulated by the **equilibrium** amongst surface **wind input**, **non-linear wave-wave interactions**, and **wave energy dissipation** [Thomson *et al.*, 2013]

Analyzed spectrum models:

- Donelan and Pierson, (1987)
- Elfouhaily *et al.*, (1997);
- Pierdicca and Pulvirenti, (2008)
- Hwang *et al.*, (2011)

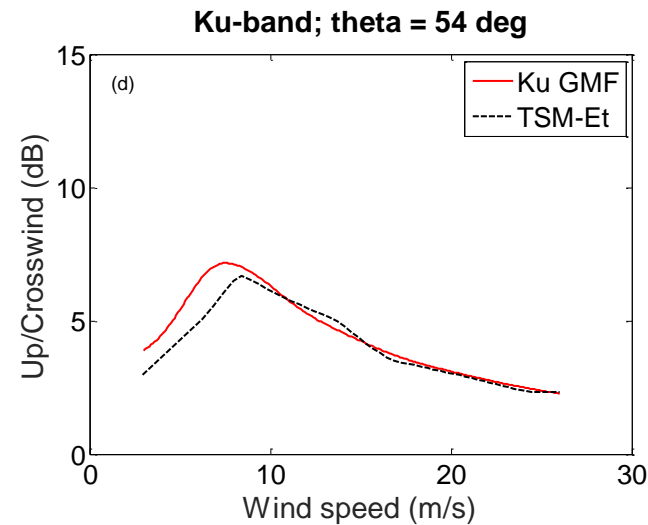
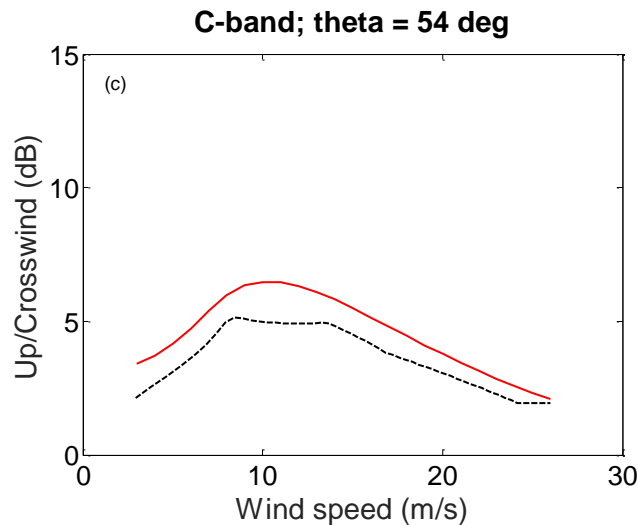
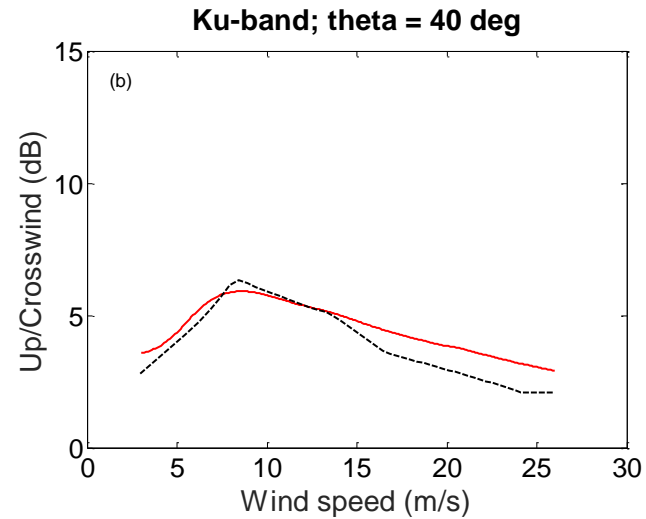
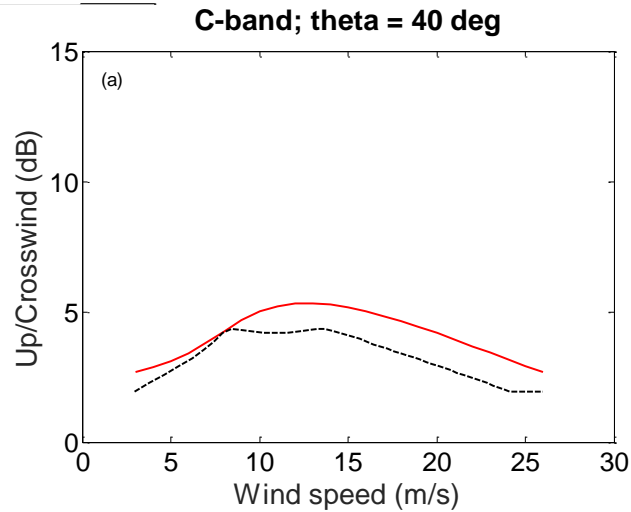


**We have decided to use the Elfouhaily spectrum model**

With a tuned spreading function, according to Pierdicca and Pulvirenti (2008).

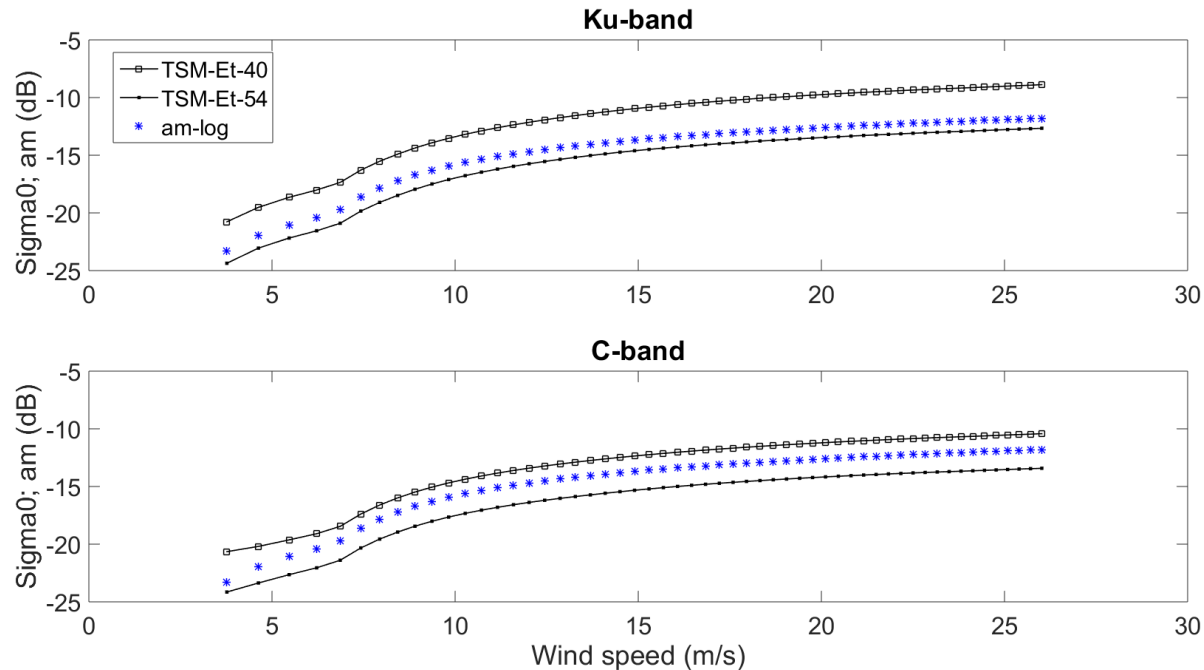
# Ocean Wave Spectrum Models (2/2)

The Et spectrum is ingested in the TSM (TSM-Et). **Disagreement** seen at **both C- and Ku-band** between **TSM-Et** and the corresponding **GMF**



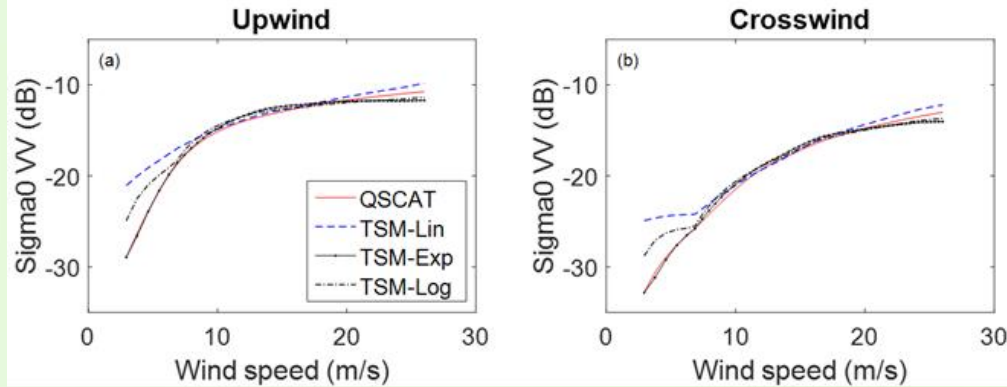
# Our Modified Ocean Wave Spectrum (1/3)

We have recognized the **equilibrium range parameter  $\alpha_m$**  of the short wave spectrum as the principal contribution in the ocean response to surface wind speed

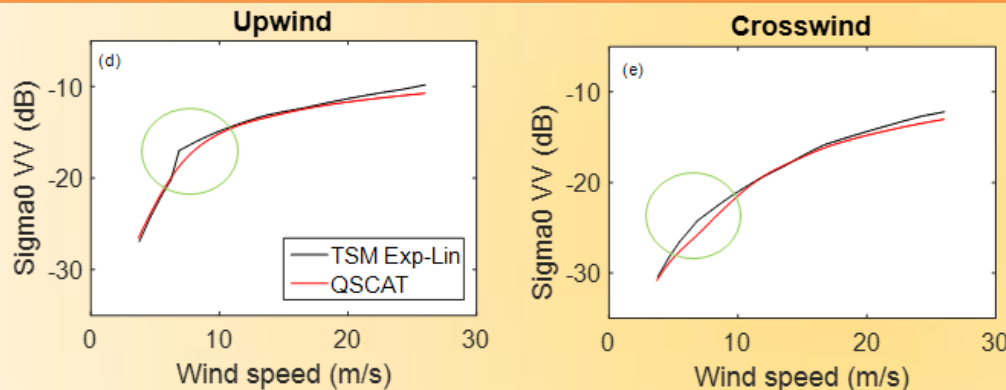


*It is a parameterization of the equilibrium condition among non-linear wave-wave interactions, wind forcing and water viscosity.*

# Our Modified Ocean Wave Spectrum (2/3)



Analysis of the TSM when different  $\alpha_m$  expressions (available in the literature) are included.



TSM-Lin  $\rightarrow$  high winds;  
TSM-Exp  $\rightarrow$  low winds

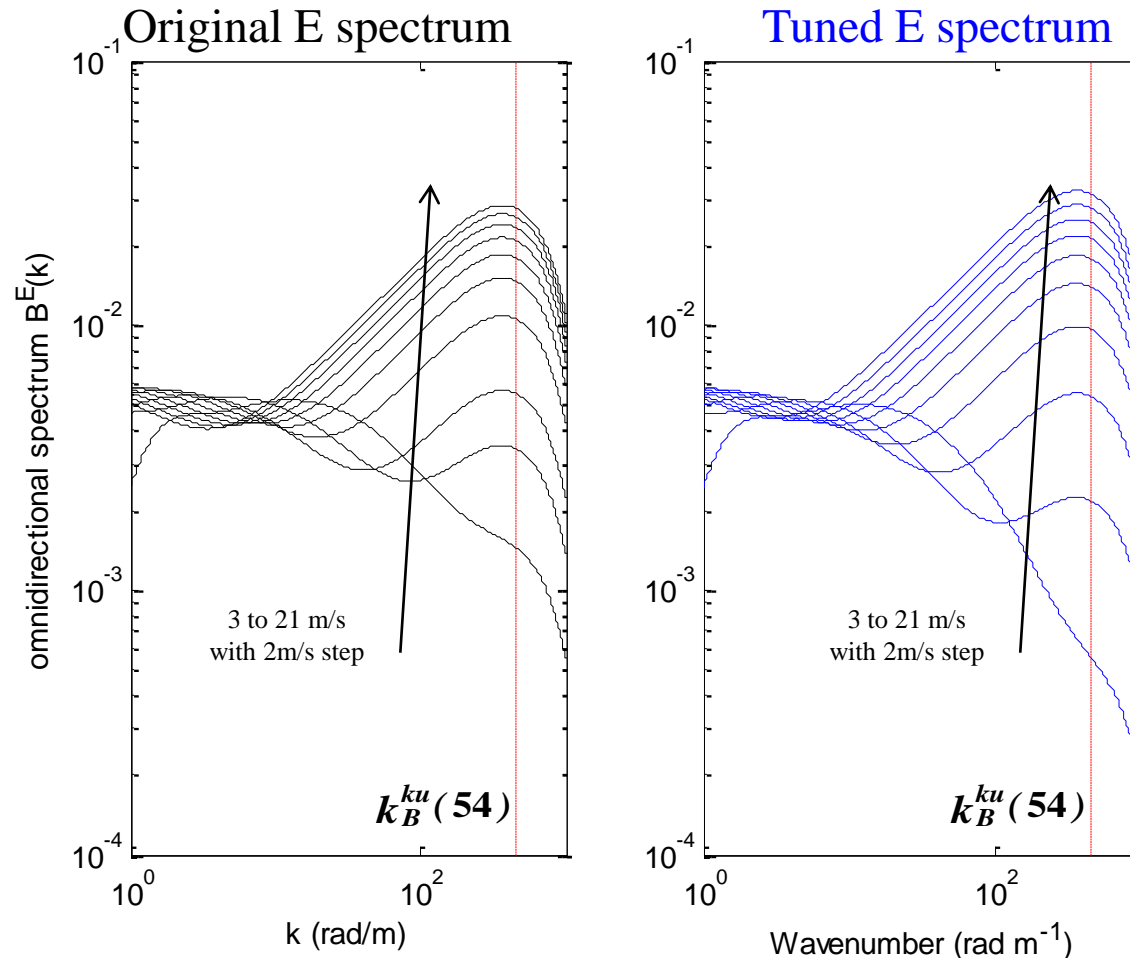
**A two-range logarithmic law as proposed by Elfouhaily *et al.* causes an inconsistency at 7 m/s**

**We have defined a new unified law**

$$\alpha_{m\_NEW} = (1-h)\alpha_{m\_exp} + h\alpha_{m\_lin}$$

**$h$ : a new patching function of the wind speed**

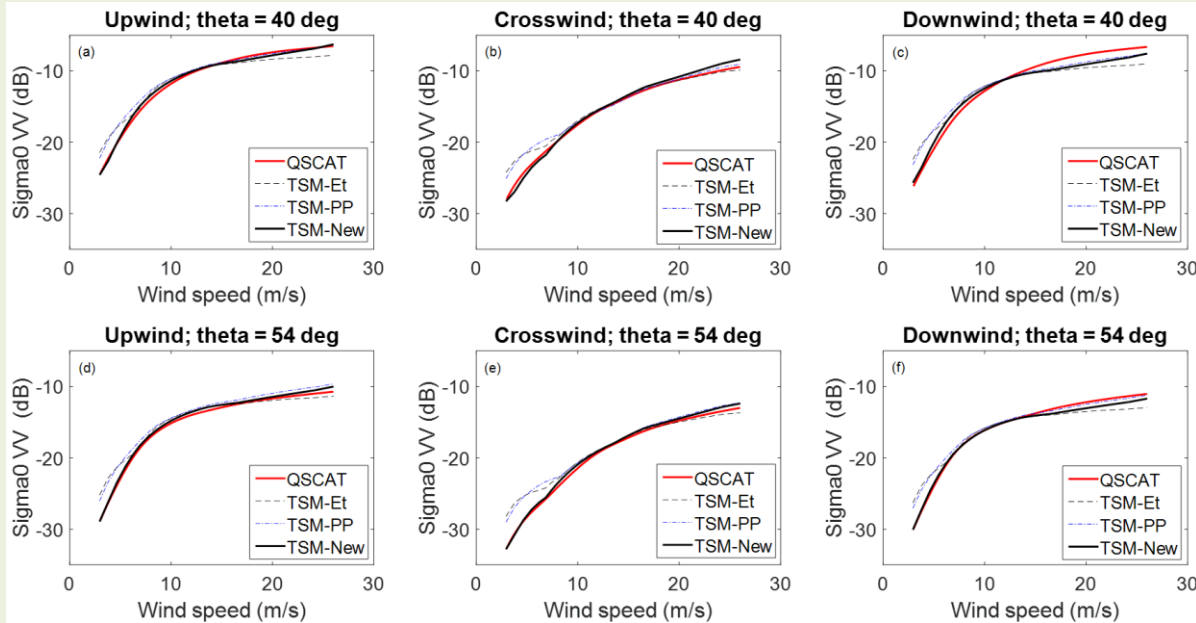
# Our Modified Ocean Wave Spectrum (3/3)



- We are basically **modifying the spectrum modulation with respect to the wind speed**. Such modulation affects the shape/sensitivity of the surface backscattering coefficient.
- Smoother transition between low and high winds

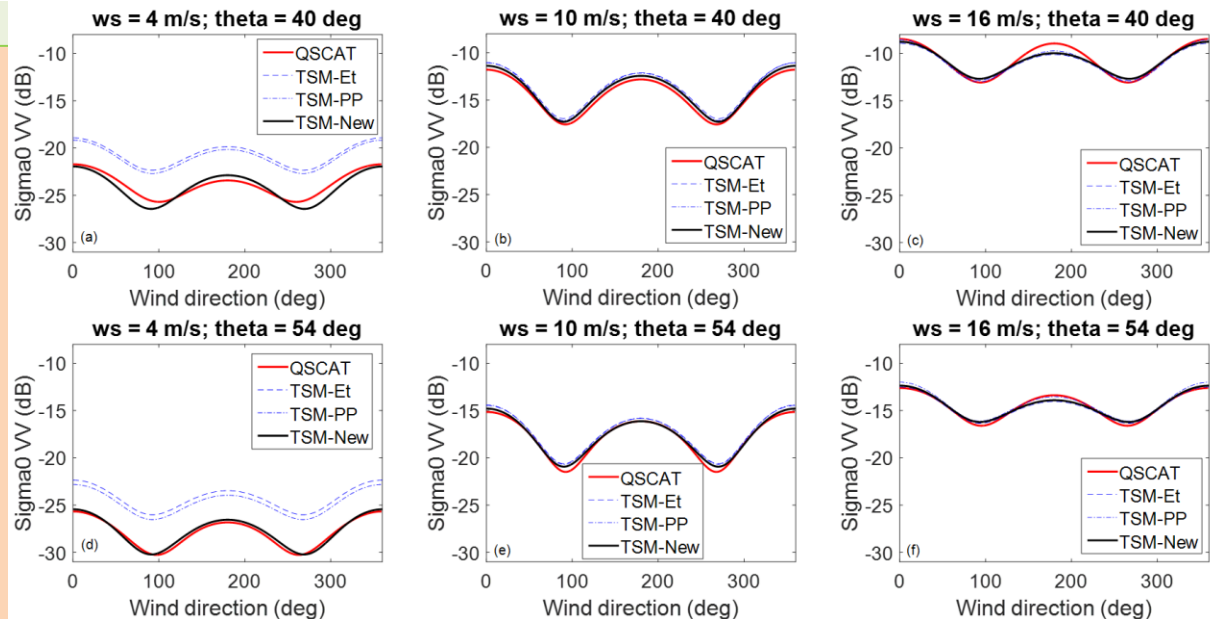
# Validation at Ku-band using GMF

$\sigma^0$  VV vs. wind speed



$\sigma^0$  VV vs. wind dir.

Improved agreement especially at low wind regimes



# Validation at Ku-band using actual data

## Dataset:

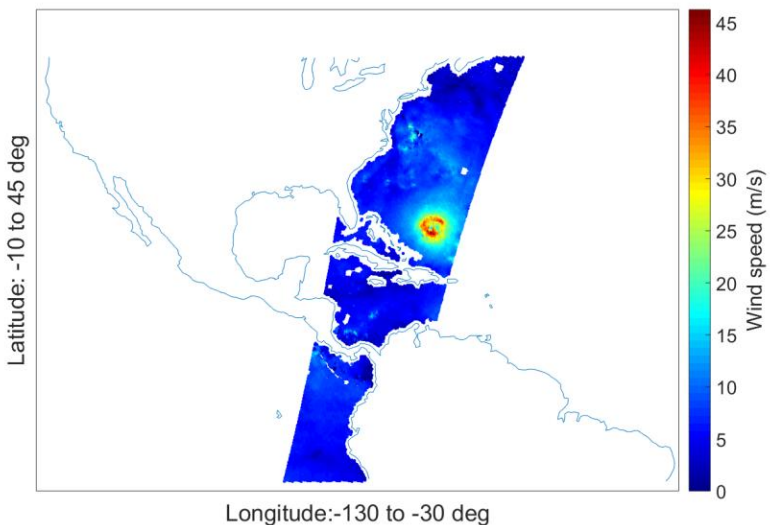
Hurricane Isabel (2003)  
NASA Ku-band (~13.4 GHz) SeaWinds Scatterometer  
Advanced Earth Observation Satellite Mission

## Analysis:

Comparison between measured  $\sigma^0$  ( $\text{sigma0}_m$ ) and  
TSM-simulated  $\sigma^0$  ( $\text{sigma0}_s$ )

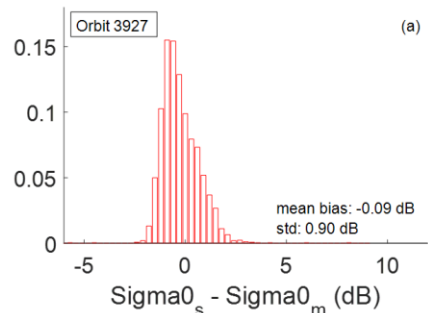
## Results:

Good agreement at Ku-band

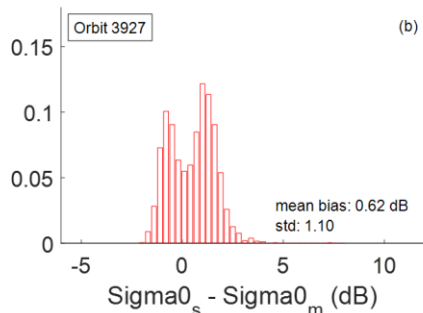


**Aft look**

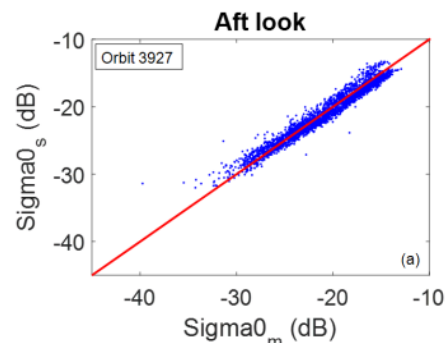
**Fore look**



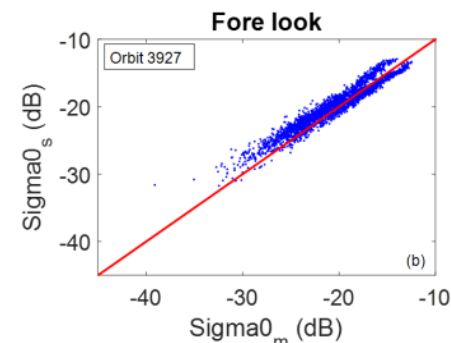
(a)



(b)



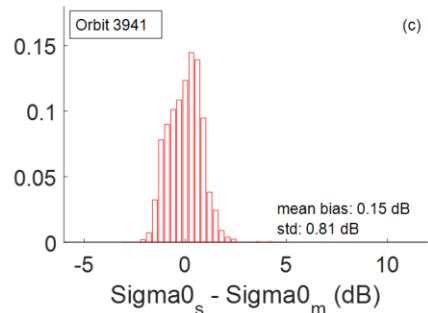
(a)



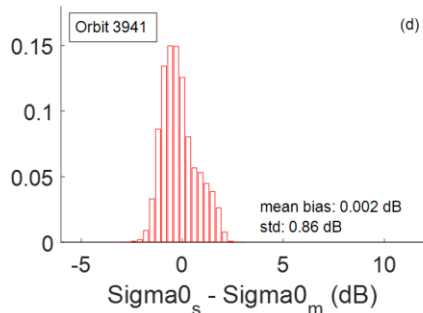
(b)

**Aft look**

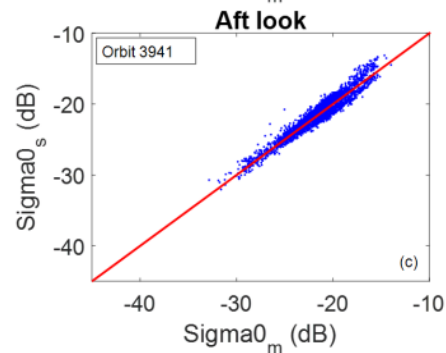
**Fore look**



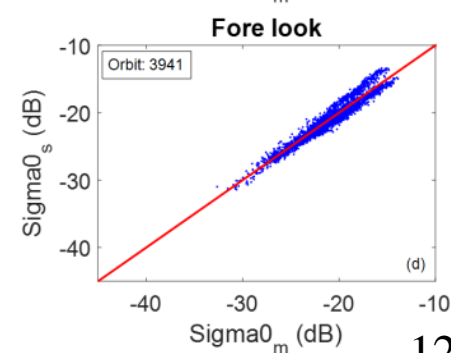
(c)



(d)

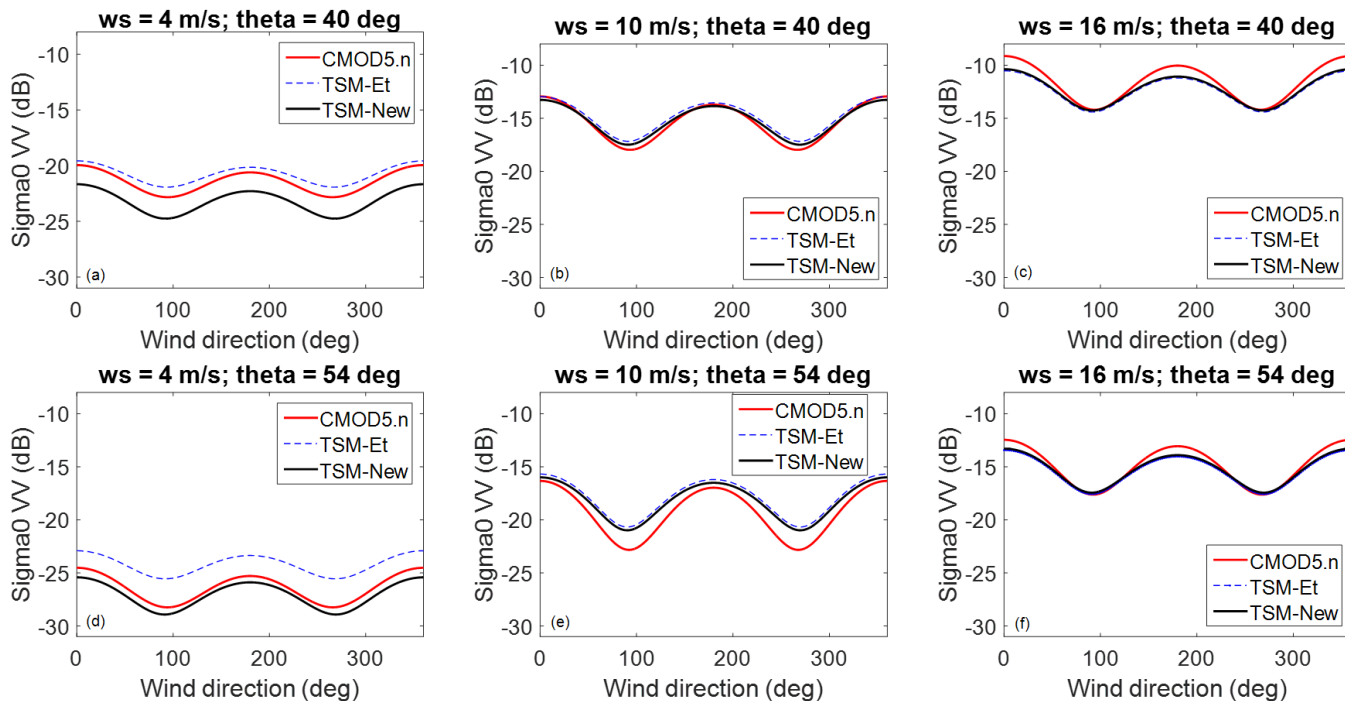


(c)



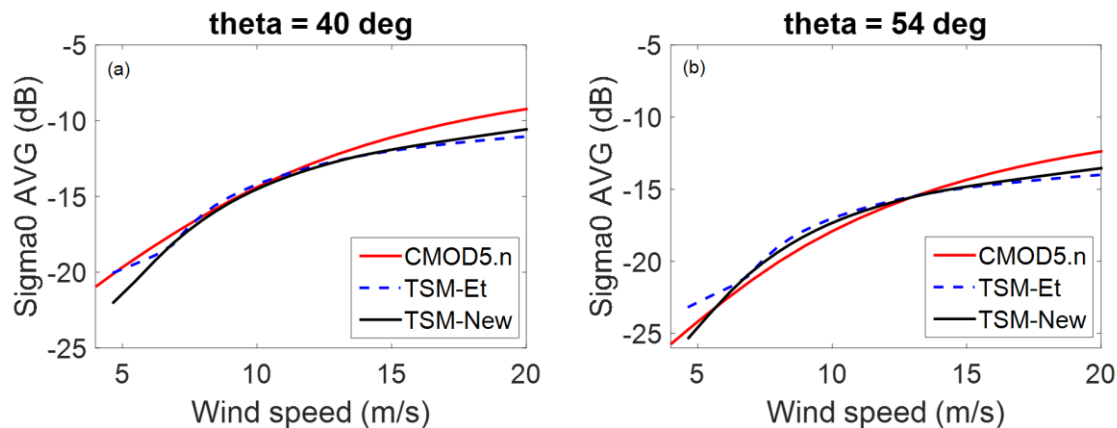
(d)

# Validation at C-band using GMF

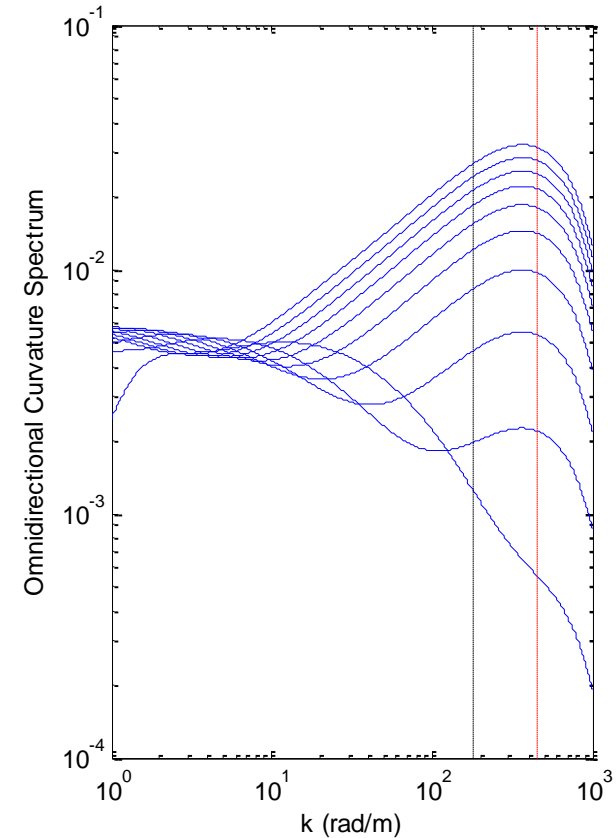
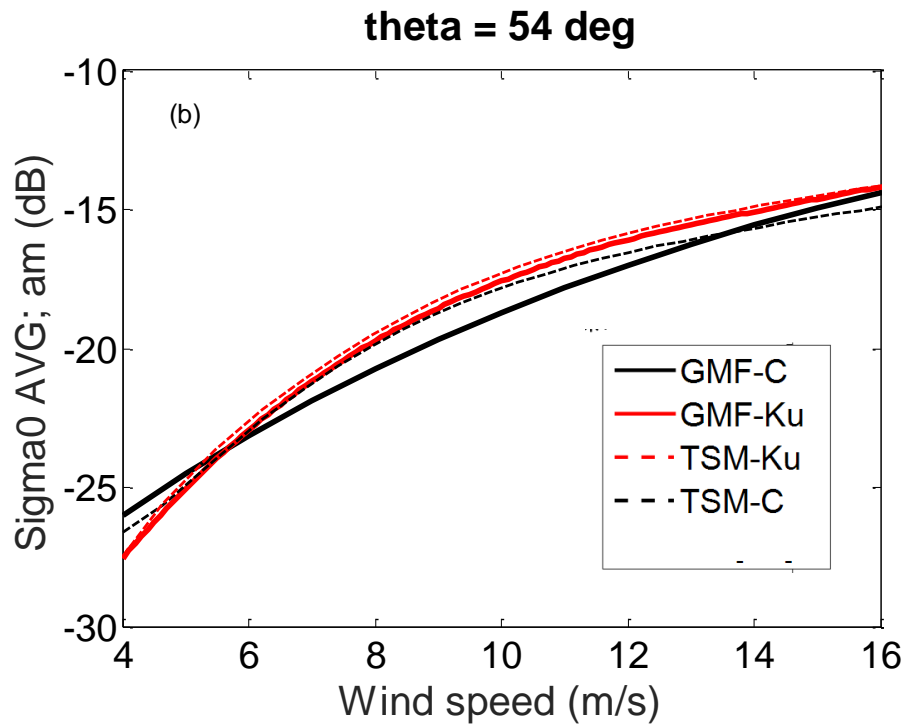


TSM-New slightly better agrees with CMOD5.n, especially for wind speeds higher than 8 m/s.

In both wind speed sensitivity and directional modulation, **the agreement in Ku-band is better than the C-band case.**



# Addressing Sensitivity to Frequency



- If we want to reproduce the sensitivity at C-band we need to **change the shape/modulation of the spectrum** with respect to the wind speed in the C-band Bragg waves range.
- This can be done by modifying the equilibrium range parameter with respect to the ocean wavenumber.
- This suggests that the **condition of equilibrium changes amongst the different ocean wave components**

# Outline

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- Theoretical Modeling in Presence of Wind
- **Theoretical Modeling in Presence of Both Wind and Rain**
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# Rain Effects: Problem Overview

## Objective:

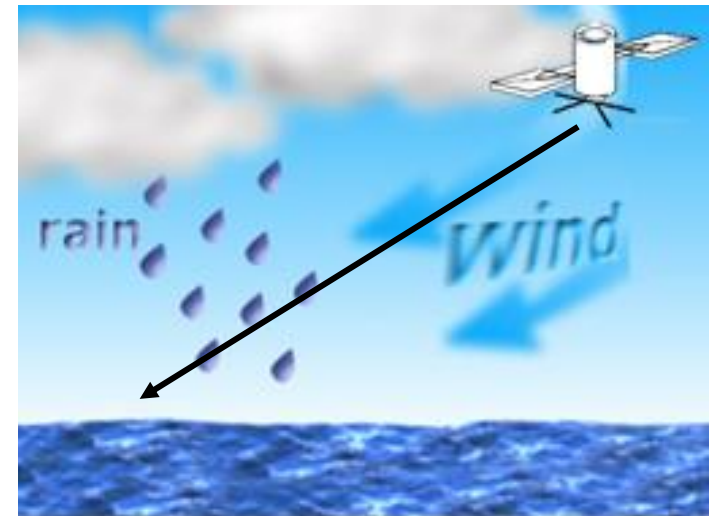
Modeling the effects of the rain on the surface backscattering coefficient.

## Methodology:

- We have focused on the surface effects.
- *The spectrum of the short waves has been modified in order to include ring waves and wave damping effects.*

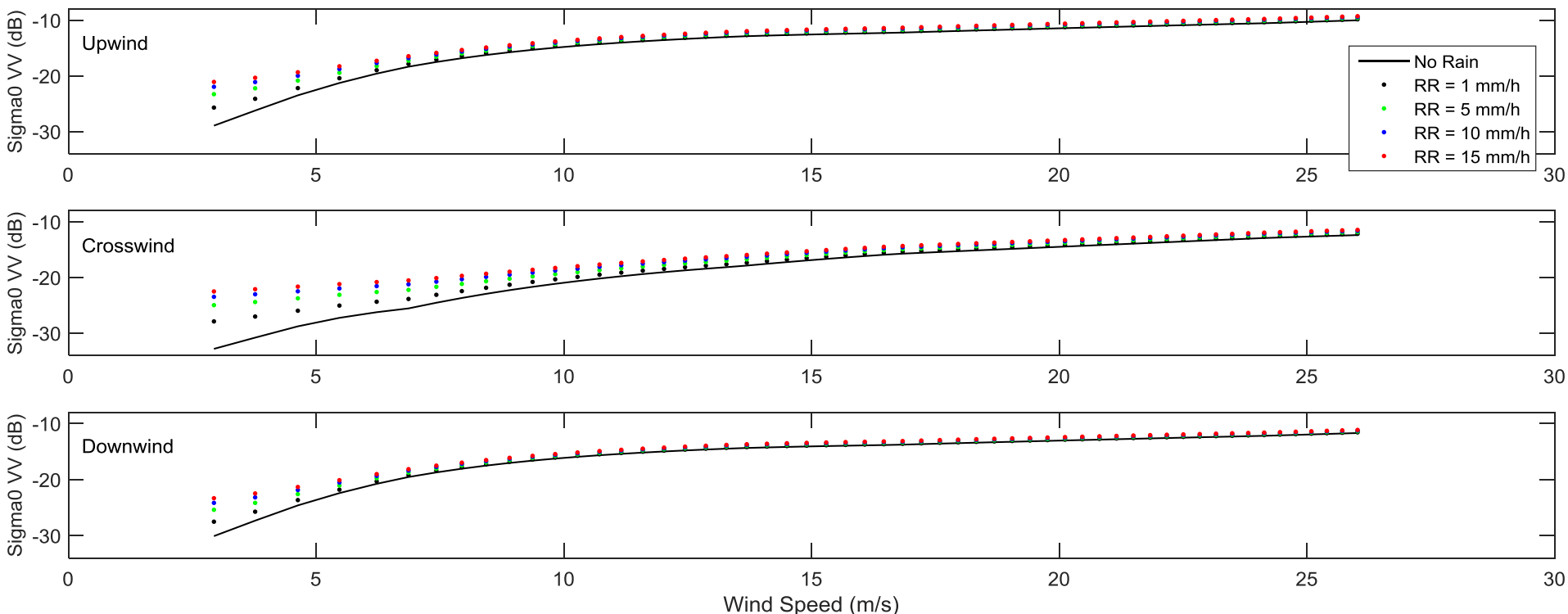
## Motivations:

- More accurate approach in estimating wind than empirical methods
- Opportunity to ingest an inversion algorithm to jointly estimate wind/rain
- Opportunity to evaluate the uncertainty of the rain rate estimates which, in turn, affect the uncertainty in the wind speed and direction retrievals



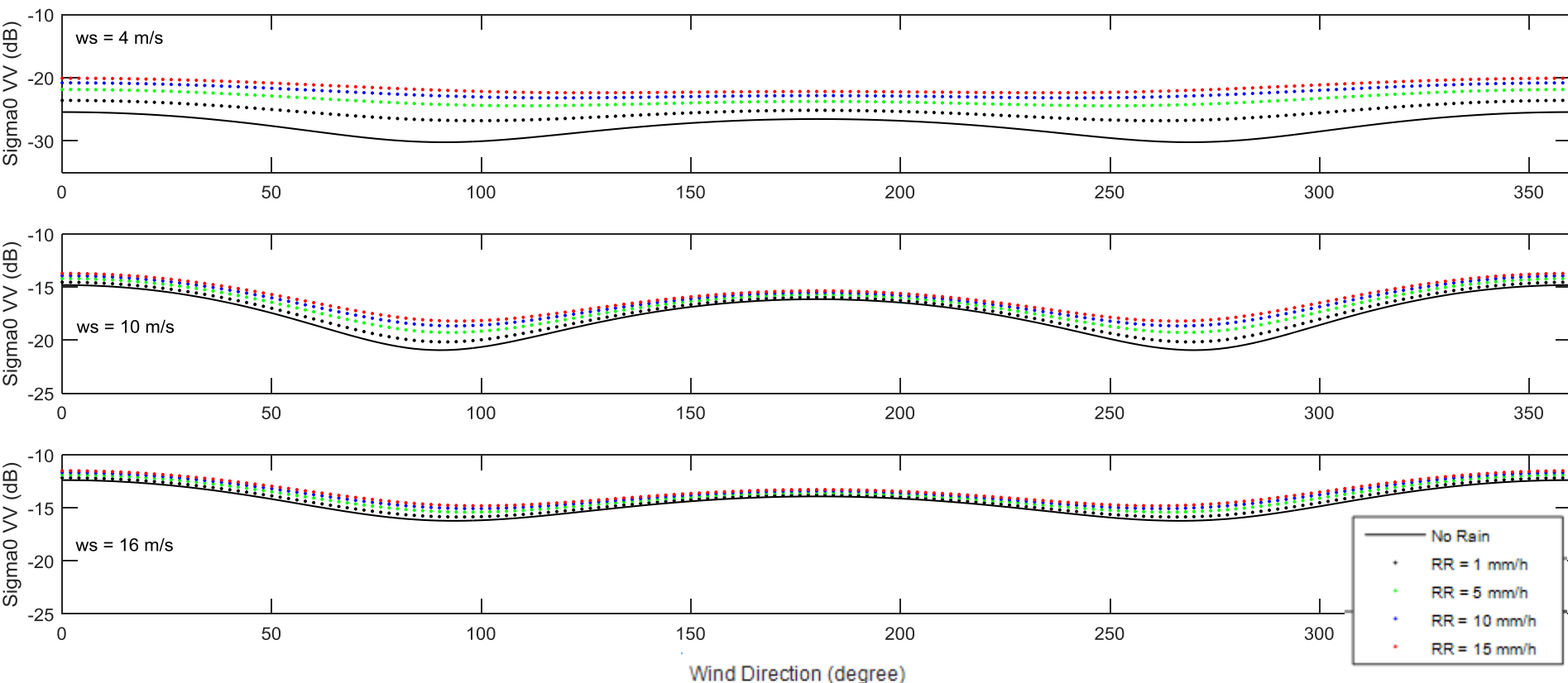
[image adapted from:  
<http://www.aquatic.uoguelph.ca/oceans/PacificOceanWeb/NPPProperties/Salinity.htm>]

# Ocean Backscattering Sensitivity to Rain



- Rain increases the magnitude of the surface backscattering  $\sigma^0$ ;
- Rain effects are more evident at low wind regime as it increases the roughness of the surface
- At higher wind regimes the wind-induced roughening effects are dominant.

# Ocean Backscattering Sensitivity to Rain

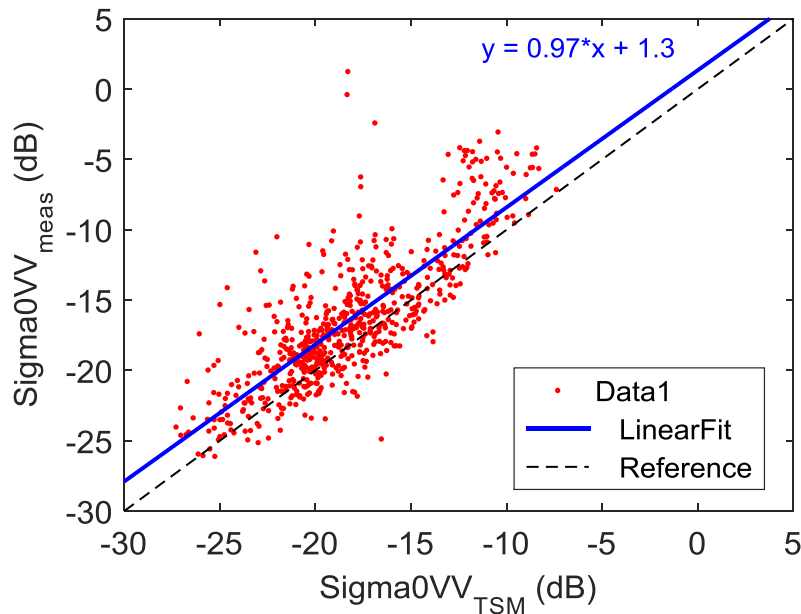


Rain reduces the upwind and crosswind asymmetry, which is essential to estimate wind direction from scatterometer observations;

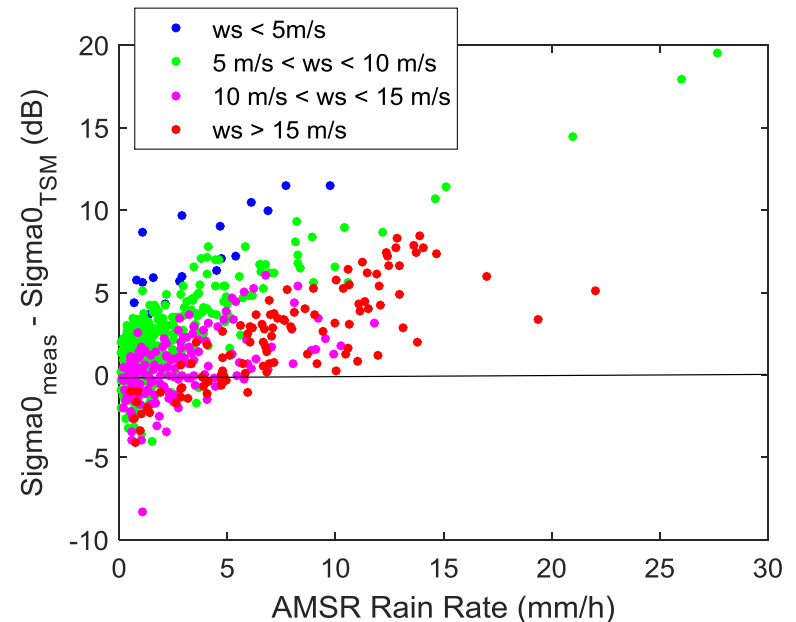
# Preliminary Results with Actual Data

Hurricane Isabel (2003)

Wind measurements: NASA Ku-band SeaWinds, Advanced Earth Observation Satellite Mission  
Rain measurements: Advanced Microwave Scanning Radiometer



The model slightly underestimates the actual data. Mean bias = 1.8 dB and std = 2.8 dB



The disagreement increases as the rain rate becomes higher.

Additional roughening contributions such as craters and stalks need to be accounted for.

# Outline

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- Science Background
- Theoretical Modeling in Presence of Wind
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- **Conclusion**

# Conclusion

## Development of a dual-frequency spectrum-based theoretical models of the ocean surface response in presence of both wind and rain

### □ Non-rainy conditions:

- Good agreement between the proposed theoretical and the Ku-band QuikSCAT GMF as well as using SeaWinds actual data.
- The agreement seen with the C-band GMF is not as good as the Ku-band case
- The model has to account for the different sensitivity to the wind in both the directional and omnidirectional ocean wave spectrum components.

### □ Rainy conditions:

- The proposed theoretical model consistently represent the rain effects on the surface backscattering
- The rain increases the ocean backscattering and it reduces the directional component
- The results with the SeaWinds/AMSR actual data are consistent to previous work [Contreras and Plant, 2006]

*This work is intended to improve the scatterometer response in support of the current empirical GMFs*

# Acknowledgements

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I really cannot find the right words to express how much grateful I am to my advisors **Dr. Joe Turk** (Jet Propulsion Laboratory) and **Dr. Svetla Hristova-Veleva** (Jet Propulsion Laboratory). Thank you for your technical help, valuable advices, your patience during our meetings and for letting me be part of the JPL ocean scatterometry group. This was a unique experience from the scientific and personal point of view. I will never forget anything of that. I would like to thank also **Dr. Ernesto Rodriguez** (Jet Propulsion Laboratory) and **Dr. Bryan Stiles** (Jet Propulsion Laboratory) for their valuable scientific support and for sharing their knowledge anytime I needed. Thanks to **Dr. Ziad Haddad** (Jet Propulsion Laboratory) for giving the chance to join his group.

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# Addressing Frequency Sensitivity

theta = 40 deg

